

# Long-term changes of the upper stratosphere as seen by Japanese rocketsondes at Ryori (39°N, 141°E)

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**Abstract.** Wind and temperature profiles measured routinely by rockets at Ryori (Japan) since 1970 are analysed to quantify interannual changes that occur in the upper stratosphere. The analysis involved using a least square fitting of the data with a multiparametric adaptative model composed of a linear combination of some functions that represent the main expected climate forcing responses of the stratosphere. These functions are seasonal cycles, solar activity changes, stratospheric optical depth induced by volcanic aerosols, equatorial wind oscillations and a possible linear trend. Step functions are also included in the analyses to take into account instrumental changes. Results reveal a small change for wind data series above 45 km when new corrections were introduced to take into account instrumental changes. However, no significant change of the mean is noted for temperature even after sondes were improved. While wind series reveal no significant trends, a significant cooling of 2.0 to 2.5 K/decade is observed in the mid upper stratosphere using this analysis method. This cooling is more than double the cooling predicted by models by a factor of more than two. In winter, it may be noted that the amplitude of the atmospheric response is enhanced. This is probably caused by the larger ozone depletion and/or by some dynamical feedback effects. In winter, cooling tends to be smaller around 40–45 km (in fact a warming trend is observed in December) as already observed in other data sets and simulated by models. Although the winter response to volcanic aerosols is in good agreement with numerical simulations, the solar signature is of the opposite sign to that expected. This is not understood, but it has already been observed with other data sets.

**Key words.** Atmospheric composition and structure (evolution of one atmosphere; pressure, density, and temperature) · Meteorology and atmospheric dynamics (middle atmosphere dynamics)

## 1 Introduction

Long-term observations of the physical parameters of the stratosphere are crucial to monitor possible climate changes. Anthropogenic temperature changes induced by increases in CO<sub>2</sub> or decreases in ozone has led to a number of modelling studies. If most of the models predict a global cooling of the stratosphere, the expected magnitude and the spatial distribution of such cooling is still being debated since the spatial and seasonal distribution of ozone trends is still uncertain. Also many feedback mechanisms, such as dynamical processes, can be expected to modify the long-term atmospheric changes. Analysis studies of the US and Soviet rocket data bases all confirm such a cooling. While recent quantifications of such trends in the stratosphere over the tropical and subtropical regions reveal a larger cooling than those estimated by numerical simulations (Angell, 1987; Keckhut *et al.*, 1999; Dunkerton *et al.*, 1998), mid-latitude estimations present smaller amplitudes (Kokin *et al.*, 1994; Keckhut *et al.*, 1995). This is probably due to a dynamical contribution that is expected to be larger in this area (Rind *et al.*, 1990). Also, while the atmospheric response to solar changes is well represented by photochemical model simulations in tropical regions (Hood *et al.*, 1993; Dunkerton *et al.*, 1998), mid-latitude responses in winter (Kokin *et al.*, 1994; Keckhut *et al.*, 1995) are not understood. Several studies have suggested a dynamical origin through feedback effects without any clear proposed mechanism. The propagation of planetary waves generated in the troposphere is affected by stratospheric circulation and in some conditions cannot propagate into the stratosphere. Also the stratosphere can determine, to some extent, its own wave driving. The upward transmissivity of wave activity as a function of temperature and wind structure have been emphasised by many investigators (Karoly and Hoskins, 1982; Chen and Robinson, 1992).

Observations of wind and temperature by rockets at Ryori (Japan) can provide a supplementary source to quantify long-term changes of the stratosphere at mid-latitudes. A previous analysis was made with temperature data extending up to 1989 (Komuro, 1989). Here we propose to update this study with a longer data set including data extended to 1995 and to add a wind analysis to the temperature. The previous analysis is improved in taking into account natural variability in using a multi-function regression model.

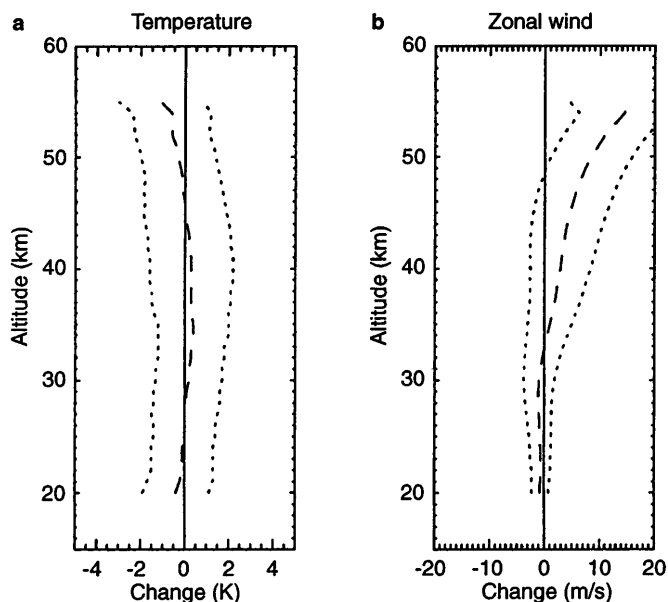
The Ryori site (39°N, 141°E) is especially useful regarding the global stratospheric circulation compared to the other subtropical rocket sites because it is located close to an anticyclone centred over the Aleutian Islands in the North Pacific Ocean at 55°N, 175°W. Maximum geopotential heights are found in the lower stratosphere over this area. This phenomenon has come to be known as the Aleutian High. A variety of observational studies have explored its structure and behaviour (Harvey and Hitchman, 1996). Quasi-stationary continental-scale circulation plays a major role in the annual cycle of the subtropical stratosphere that affects the southern winter flow. Stratospheric anticyclones have been found to be an essential aspect for air mixing and for the advection from the tropics to higher latitudes (McIntyre and Palmer, 1983; Randel *et al.*, 1993) but also for planetary wave propagation and breaking (Dunkerton and Delisi, 1986; O'Neill *et al.*, 1994). In this study a particular emphasis is placed on winter data when North Pole ozone depletion is observed and when planetary waves can propagate into the stratosphere and may contribute significantly to interannual changes and climate feedback. It is not possible to draw firm conclusions using data from a single location. However, in the stratosphere (and mostly in the upper stratosphere) few continuous data bases are available on a long-term basis for studying interannual variability and many of them require specific analysis to clean-up data pollution by instrumental effects. Also, comparisons with published multi-dimensional simulations permit interpretation of some of the observed features. This study is proposed to continue the global efforts of documenting interannual change using existing long data series as a basis for comparison with models and for theoretical interpretation.

## 2 Data analysis procedure

Meteorological rocket sounding has been in use for routine measurements of wind and temperature in the upper stratosphere at Ryori (39°N, 142°E) for many years. The instruments consist of a nose cone containing a parachute, and a temperature sensor with its transmitter, that is separated from the rocket motor at the apogee. The payload known as "Echosonde" is a pulse-position-modulation system containing a nickel-iron wire as the temperature sensing device. Temperature is measured during descent under a slink parachute while telemetering information to ground. The altitude of the sonde and the wind measurements is deduced from the

position of the sonde determined by a ground tracking radar. These rockets reach a maximum height of 62 km while the sensor is deployed at 60 km limiting measurements to the upper stratosphere, a lower altitude range compared to US and Soviet vehicles. Those measurements have been carried out and archived by the Japan Meteorological Agency. Rockets have been launched at 2 UT with no deviation larger than 30 min. This is a guarantee against any interference with atmospheric tides (Keckhut *et al.*, 1996) or diurnal instrumental heating (Hoxit and Henry, 1973) which may exist in the US rocket data base (Keckhut *et al.*, 1999). The rocketsonde vehicle, is a MT-135 rocket, a solid-propellant single stage rocket that has not been changed since July 1970. Data have been corrected for aerodynamic, solar radiation, and Joule heating, and in a lesser extent for long wave radiation and lag errors (Yata, 1970). The total temperature correction at 60 km is approximately 22°. While the sensor type has not been changed since 1970, the thermistor mount was modified in 1977. The rocketsonde ES89P actually in operation has replaced the ES-64B sonde previously used. The change concerns mainly the wire frame size. Since April 1975, wind data are also corrected for fall velocity and ballistic motion effects. Such changes are probably minor and smaller than instrumental noise. However they may have induced some changes of the mean which need to be considered for long-term data analysis. Two to four launches per month were made systematically since 1971 providing a total number of vertical temperature and wind profiles of 794. The confidence for trend estimates depends on the correlation between successive measurements, the atmospheric variability, and the sampling. Usually, in the stratosphere the lowest uncertainty is obtained for data averaged over 15 to 30 days. In this study, it was decided to use monthly means. The number of monthly means obtained with the whole data set of Ryori is 277. The soundings are uniformly distributed as a function of time and the number of monthly means is around 69–70 for data partitioning within any of the three month intervals and 23–24 when any single month analysis is selected.

The data analysis for trend detection used a least square fitting of the data with a multiparametric adaptative model that has been developed to take into account simultaneously natural variability and possible trend. The AMOUNTS (adaptative model for unambiguous trend survey) regression analysis has already been used and described in previous lidar (Hauchecorne *et al.*, 1991; Keckhut *et al.*, 1996) and rocket (Keckhut *et al.*, 1999) trend studies. The regression model used here includes a seasonal trend, a seasonal linear function related to solar activity (10.7 cm solar flux), a seasonal function of the quasi biennial oscillation, a seasonal function of the optical depth induced by volcanic stratospheric aerosols (Sato *et al.*, 1993), and the seasonal components (annual and semi-annual). This model also includes step functions at some specific dates when sensor mount or correction applied have been changed (as mentioned previously). The largest bias detected occurs for wind above 45 km when new



**Fig. 1a, b.** Amplitude of step functions (*dashed lines*) as a function of altitude in **a** the temperature and **b** wind data series computed with the regression model respectively in January 1977 and April 1975. *Dotted lines* are the  $2\sigma$  confidence intervals. Changes are shown for temperatures in K and in  $\text{ms}^{-1}$  for wind

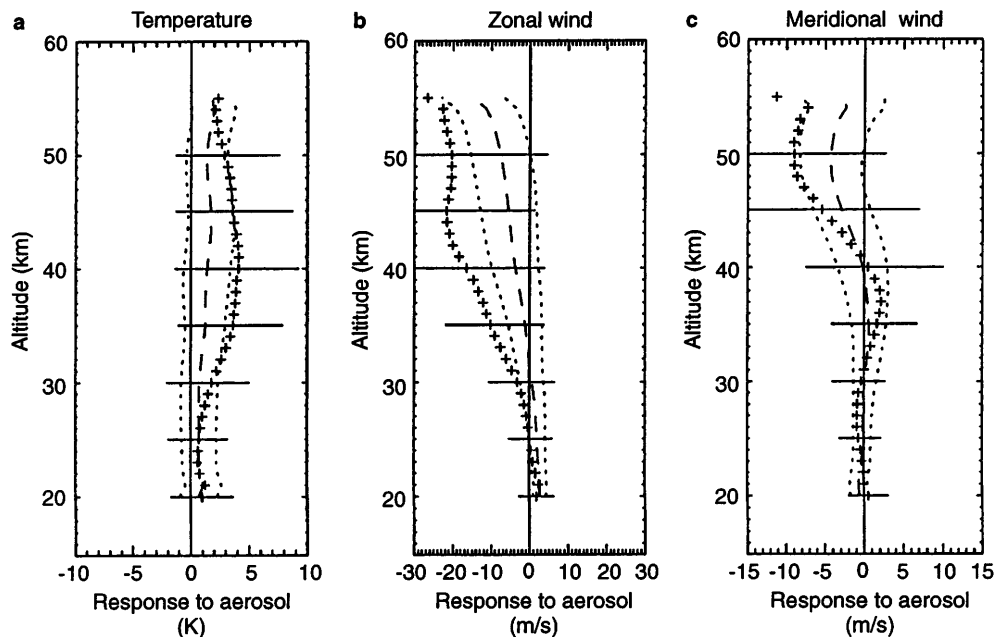
corrections were introduced in the data analysis (Fig. 1b). On the other hand, no significant change of the mean is noted for temperature (Fig. 1a) after the sondes were improved.

### 3 Results of long-term changes

#### 3.1 Volcanic eruptions

Volcanic eruptions may have a significant impact on the mean atmospheric temperature. Direct effects in the

lower stratosphere and troposphere over tropical regions are unambiguously detected after major volcanic events such as the last Mount Pinatubo eruption (Labitzke and McCormick, 1992; Angell, 1997). Indirect effects, caused by the vertical radiative balance and effects on the global circulation, may occur elsewhere, even far from the region of the direct forcing (Rind *et al.*, 1992). Depending on the occurrence of such events in the data series, trend analysis may be influenced and miss-estimated. During the period analysed here, three major eruptions have injected particles into the stratosphere. If it is assumed that volcanic aerosols induce climatic perturbations linearly related to the optical depth, the computed mean annual response to stratospheric volcanic aerosols found over Ryori is not highly significant and is between 1 and 2 K for temperature for a mean optical depth of 0.1 (Fig. 2a). This result is in agreement with estimations made after the El Chichon eruption by Angell (1997) from US rocketsonde. The amplitude of the temperature response in the upper stratosphere is predicted to be minimum at mid-latitudes by some model simulations of the forcing of an uniform stratospheric aerosols loading performed with the GISS (Goddard Institute for Space Studies) global climate model (Rind *et al.*, 1992). A larger mean annual response is estimated here. Winter response presents an enhanced warming (3–4 K) in good agreement with the signature reported above the Observatory of Haute-Provence in France (Keckhut *et al.*, 1995). At high latitude, in the upper stratosphere and lower mesosphere, simulations show that temperatures warm in winter, with a magnitude from one to four times the model's natural variability. This arises due to changes in dynamics, in particular the altered residual circulation. A similar warming is observed using a spectral three-dimensional model of the stratosphere (Pitari, 1992) revealing a zonal mean warming of 2 K extended down to  $45^\circ$  induced by the modification of the propagation



**Fig. 2a–c.** Vertical response in **a** temperature, **b** zonal wind and **c** meridional wind series (in K and in  $\text{ms}^{-1}$ ) to a mean optical depth change of 0.1, induced by stratospheric aerosols. *Dashed lines* are the analyses of all the data, while *dotted lines* are the  $2\sigma$  confidence intervals. Results for winter month (December–January–February) analyses are shown by *crosses*. *Horizontal lines* are associated with confidence intervals

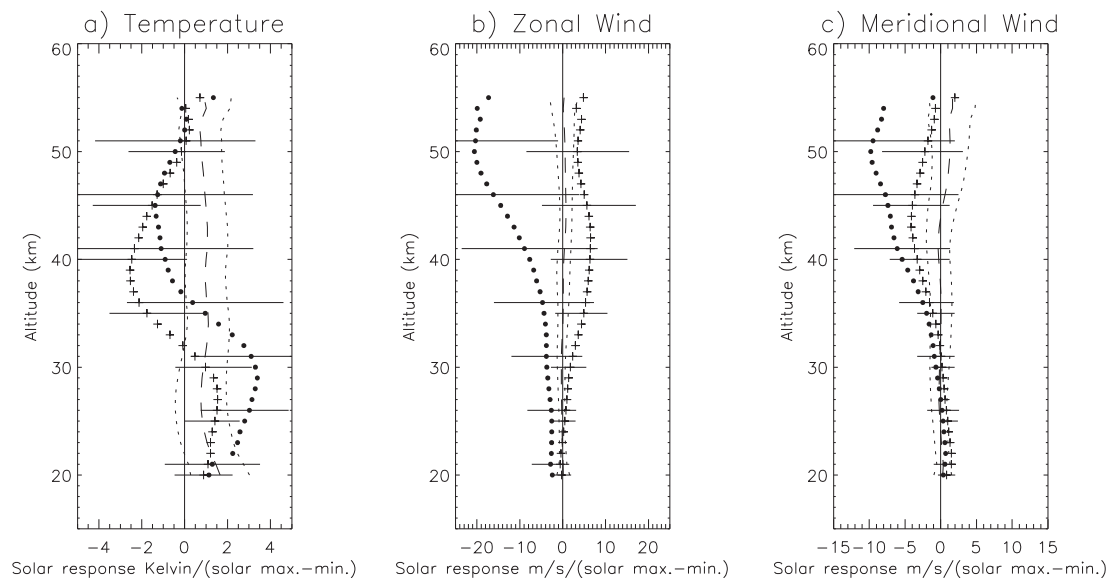
and the energy generation in the troposphere of planetary waves through changes of the vertical static stability at low latitudes. However, the response is not uniform with longitude and a region of maximum is expected over eastern Asia and to a lesser extent over North America (Pitari, 1992). Models and observations both present a maximum zonal wind change around the stratopause (Fig. 2b). Zonal wind changes in the model are smaller ( $8 \text{ ms}^{-1}$ ) than the ones reported here ( $20 \text{ ms}^{-1}$ ) but may also be associated to the regional warming. A nearly significant meridional anomaly is also observed around the stratopause (Fig. 2c).

### 3.2 Solar activity

Temperature and ozone are expected to be correlated with solar activity with a maximum around the stratopause following the line of zero zenith angle. This direct radiative and photochemical effect is calculated selecting all the six summer months of data from April to September (Fig. 3). The temperature response is quite uniform from 30 km to the top (55 km) and is equal to around 1 K, from solar minimum to solar maximum conditions, just above the confidence limit (95%). Wind responses are less significant. Zonal wind exhibits a maximum of  $0.8 \text{ ms}^{-1}$  at 45 km with a uncertainty of  $1.5 \text{ ms}^{-1}$  (95%). Temperature and wind responses, observed in this 25-y monthly rocketsonde data, to solar activity changes, are small. However, amplitude of changes for both temperature and wind associated with solar change activity, deduced from the multi-function regression analysis applied here on Japanese rocketsonde series are in good agreement with the numerical simulations of Garcia *et al.* (1984). Temperatures are expected to be correlated with solar activity with a

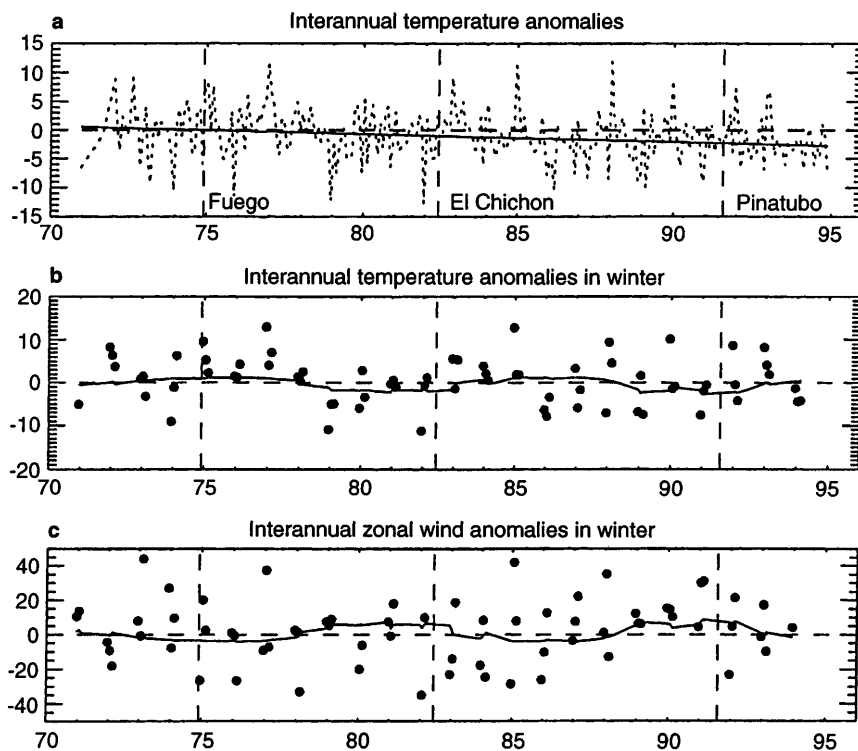
maximum of 2 K per solar cycle in tropical regions around 45 km. At higher latitudes simulations show that the temperature response is reduced with a flat response over a larger altitude range. Zonal winds should vary with an amplitude of  $1\text{--}2 \text{ ms}^{-1}$  around the mid-latitude stratopause from solar minimum to solar maximum conditions. More recent simulations (Huang and Brasseur, 1993) including more realistic solar forcing, gave a slightly smaller response. Observed solar effects on ozone in the latitude band between  $30$  to  $50^\circ\text{N}$  (Miller *et al.*, 1996) are maximum at  $35\text{--}40$  km with a magnitude of around 4–5% changes from solar maximum to solar minimum conditions in good agreement with the solar effects expected from theory. The magnitude of the observed temperature and wind response can be considered in as good agreement with theory.

During winter season, an enhanced effect, nearly significant, can be observed around 40 km on both wind ( $6.5 \pm 10 \text{ m/s}$ ) and temperature changes ( $-2.5 \pm 2.6 \text{ K}$ ) having an opposite sign (Fig. 4b, c). The analysis of December zonal wind at Ryori made by Kodera and Yamazaki (1990), using a shorter data set, reveals even larger amplitude ( $20 \text{ m/s}$ ). This unpredicted response which can not be understood with radiative and photochemistry processes alone was already observed at mid-latitude with lidar (Hauchecorne *et al.*, 1991; Keckhut *et al.*, 1995) and with Soviet rocketsonde data series (Kokin *et al.*, 1994). This seems to imply that some feedback mechanisms may be involved. Several authors have suggested an indirect role through the dynamics of the middle atmosphere and planetary waves are highly suspected (Geller and Alpert, 1980; Hines, 1974). However, some investigations (Kodera *et al.*, 1990; Balachandra and Rind, 1995) have shown that a very large and unrealistic change of the solar heating rate in the stratosphere is required to produce such a



**Fig. 3a–c.** Response to solar activity changes from solar maximum to solar minimum conditions on **a** temperature, **b** zonal wind and **c** meridional wind series (in K and in  $\text{ms}^{-1}$ ). Dashed lines are analyses of the data obtained from April to September, while dotted lines are the

$2\sigma$  confidence intervals. Results for winter month (December–January–February) analyses are shown by crosses. Results for January monthly analyses are noted by circles. Horizontal lines are confidence intervals



**Fig. 4.** Data series of the anomalies (data minus seasonal changes) at 40 km over Ryori for temperature **a** for all the data, **b** for winter temperature only and **c** for winter zonal wind. Solid lines in **b** show linear trend or **b**, **c** solar cycle response computed by the regression model

change. Simulations with a two-dimensional model including an interactive calculation of the dynamic (Huang and Brasseur, 1993), shows that feedback effects due to wave-mean flow interactions by gravity and planetary waves, seem to have small impacts on temperature and tend to reduce wind perturbations rather than to amplify it. However, 2 D model are not well adapted to study the dynamical feedback in a realistic manner. Moreover, tropical dynamics, responsible for the QBO and semi-annual oscillation (SAO), are not included in these models, and possible feedback due to small-scale dynamics is not yet fully understood and quantified.

For January, a very different correlation between solar activity and upper stratospheric wind is observed (Fig. 3). This can be due to the anticyclonic cell (the Aleutian High) developed over North Pacific which usually becomes very intense in January. If the intensity of the cell is modulated according to changes of solar activity, which induce small changes in radiative forcing, the wind speed can be modulated too. Since Japan is located not far from the northwest side of the cell, one can understand the large negative correlation between solar activity and both components of the wind (Fig. 5). However, this needs to be specifically tested with a three-dimensional model.

### 3.3 Trends

Taking into account some of the components of the natural variability described previously, a significant cooling of  $2.0 \pm 1.9$  to  $2.5 \pm 1.1$  K/decade is computed (Fig. 6), while wind series reveal no significant change. Zonal and meridional wind changes are respec-

tively smaller than  $0.6 \pm 1$   $\text{ms}^{-1}$  and  $0.1 \pm 0.5$   $\text{ms}^{-1}$ /decade. This study confirms the general cooling (Fig. 4a) of the upper stratosphere already observed with other data bases and with models. While very similar amplitudes are obtained in the upper stratosphere with the preliminary analysis done by Komuro (1989), a larger amplitude by a factor of 3 can be noted around 20–30 km. Trend results are in good agreement with recent rocket analyses (Keckhut *et al.*, 1999; Dunkerton *et al.*, 1998) that have reported a cooling around 1.8 K per decade in the altitude range between 35 to 50 km, for stations located mostly at low latitudes. However, trends in the lower stratosphere are twice those observed in tropical and subtropical regions. Also, one can note, that a lidar at a mid-latitude site in the south of France ( $44^\circ\text{N}$ ) has observed a non-significant and smaller cooling in the mid-upper stratosphere (Keckhut *et al.*, 1995) and a slight warming in winter (Hauchecorne *et al.*, 1991).

There are still some uncertainties about the spatial distribution of the ozone decrease in the lower stratosphere. Several Dobson spectrometers have been operated in Japan since 1958. These measurements seem to reveal smaller and not significant ozone decreases compared with the ones observed over Europe and North America over the last decades (WMO, 1988; Reinsel *et al.*, 1994; SPARC, 1998). Moreover, these instruments operating since the seventies in the Umkehr mode provide one of the longest record of stratospheric ozone vertical distributions. These measurements in the lower stratosphere are polluted by volcanic aerosols that require corrections (Newchurch and Cunnold, 1994) limiting the length of the data set. Also the Japanese instruments have known several interventions which reduce the confidence in the vertical profile of trend

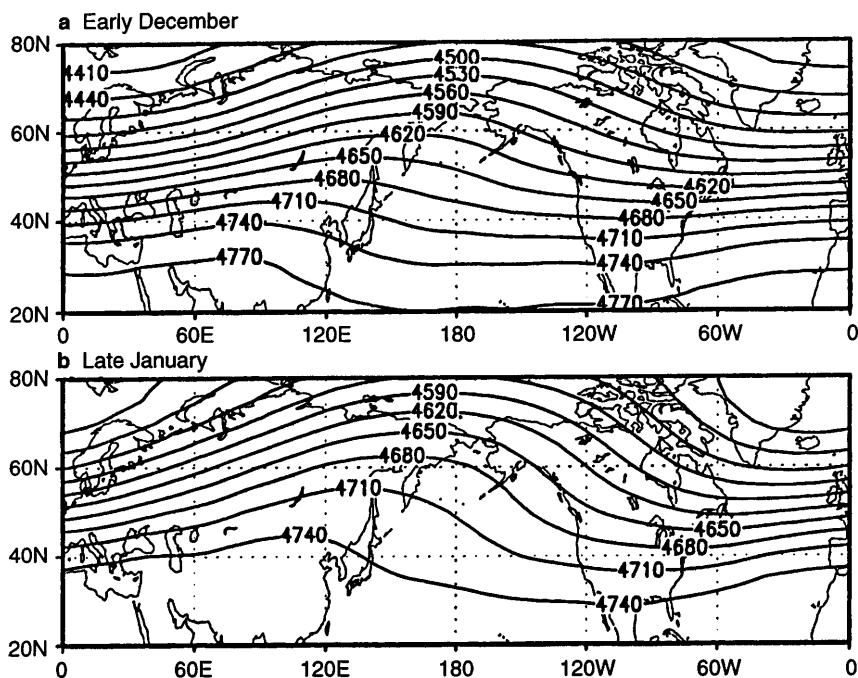


Fig. 5a, b. Climatology of the geopotential height at 1 hPa (1980–1994) for a) early December and b) late January. Units are in decameters

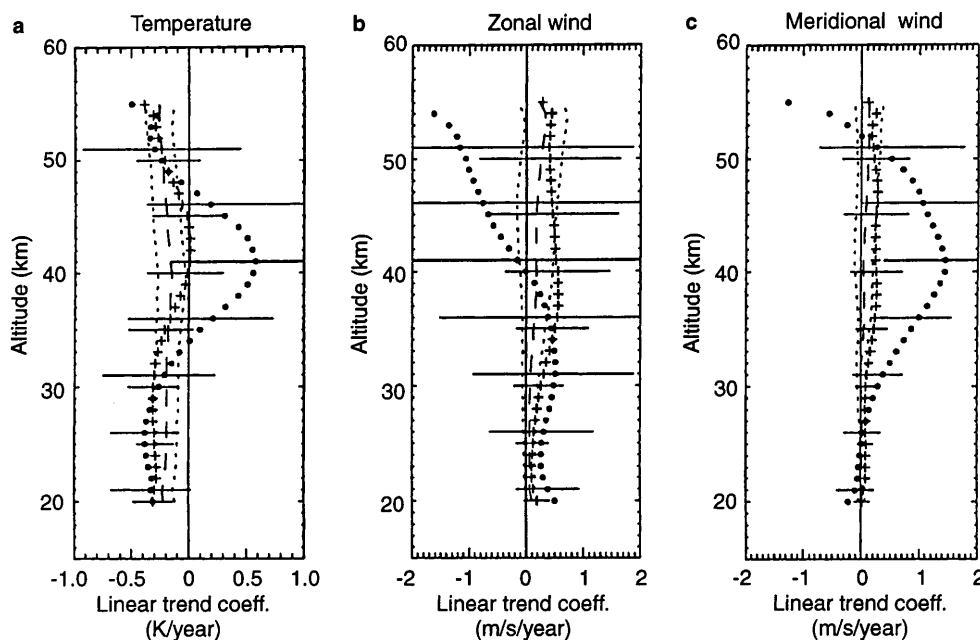


Fig. 6a–c. As Fig. 2 except for trend. In this case, circles are associated with December monthly analyses

estimates mainly in the lower stratosphere. In the latitude band between 30 to 50°N mean trend coefficients present two maxima of 7% at 45 km and 15 km. There are more uncertainties at this last altitude due to the presence of stratospheric aerosols and the atmospheric variability. Using a radiative transfer model including such an estimated ozone change and CO<sub>2</sub> increase on the period between 1970 and 1986 (WMO, 1989). Miller *et al.* (1992) have reported a cooling of 0.4 K/decade around 30 km to 1.1 K/decade at 45 km. Observations at Ryori reveal larger cooling by factors of 2 in the upper stratosphere to 4 in the mid and lower stratosphere.

Seasonal components of computed trends at Ryori are mainly caused by winter season behaviour. The winter temperature trend profile (Fig. 6a) appears to be significantly different compared to other seasons; enhanced below 35 km and above 50 km and reduced around 40–45 km. A significant enhancement in the lower stratosphere in winter ( $3.1 \pm 2.1$  K/decade) is expected as the largest ozone decline occurs in the winter-spring season as already observed (SPARC/IOCGAW, 1998) and simulated (Ramaswamy *et al.*, 1996). This last study reports an associated cooling of 1.5 K at mid-latitude which is half the one observed here. A coupled analysis of TOMS (total ozone mapping spec-

trometer) and MSU (microwave sounding unit) satellite data (Randel and Cobb, 1994) in winter (January–February) reveals some coherent horizontal patterns of ozone and temperature changes. Trend maps show a pronounced zonal wave 1 pattern, with large negative ozone and temperature trends over eastern Asia. Trends of  $-30$  DU/decade and  $-2$  K/decade are estimated in this area in quite good agreement with our results. Although no seasonal changes of trends have been reported on the recent rocket analyses at low latitudes in the upper stratosphere, the vertical winter trend pattern observed over Ryori has been already observed from other data sets at mid-latitudes from Soviet rockets (Kokin *et al.*, 1994) and from lidar (Hauchecorne *et al.*, 1991); even warming can be noted during winter period in these data sets. This signature is expected from 3 D models covering the whole middle atmosphere up to the upper mesosphere (Rind *et al.*, 1990). This is also the case for winter months over Ryori and more visible in December when a nearly significant warming of  $5$  K/decade can be noted. During this month the largest zonal wind amplitude is usually observed. Some feedback mechanisms through the dynamical processes can be expected. In this study no significant zonal wind changes (Fig. 6b) have been noted for December, while an enhanced wind trend is observed in winter months of  $5.6 \pm 8$   $\text{ms}^{-1}$ /decade. Expected trends on stratospheric zonal wind are smaller than  $1$   $\text{ms}^{-1}$  per decade for a doubling of  $\text{CO}_2$  concentration (Rind *et al.*, 1990) and predict a zero line of wind change in the vicinity of  $40^\circ$  latitude. However, published simulations of long-term stratospheric wind changes are not numerous and uncertain and few estimations of stratospheric zonal wind changes have been computed for other climate forcing such as the one associated with a non-uniform ozone decrease.

#### 4 Summary and conclusion

Temperature and wind data series obtained by successive rocketsonde launches from Japan, have been analysed to quantify possible anthropogenic trends while taking account of natural variability and possible effects of instrumental changes. While a bias is detected in the wind data series mainly above  $45$  km probably induced by corrections introduced in the data analysis, no changes associated with instrumental changes are found for temperature. Annual mean atmospheric responses to natural variability such as volcanic eruptions and the solar cycle are small and in quite good agreement with models. Although the winter response to volcanic aerosols is in agreement with numerical simulations, the solar signature is of the opposite sign compared to that expected. This signature was already observed with other data sets (Mohanakumar, 1985; Hauchecorne *et al.*, 1991) and is not yet fully understood. The significant cooling of  $2.0$  to  $2.5$  K/decade obtained in this study is larger than the one predicted by models suggesting the contribution of some feedback or some other climatic forcing. The winter cooling tends to be minimum around  $40$ – $45$  km (it even tends to warm

for December period) as already observed in other data sets and simulated by models. However, the wind series does not reveal very significant changes. Longer data series are required and the natural variability needs to be smaller. This can be achieved by obtaining more frequent measurements or providing a better fitting function of the variability. In winter, it may be noted that the amplitude of the atmospheric response is enhanced in the lower stratosphere, probably because the largest ozone depletions occur during this season. The warming effect in the upper stratosphere in winter is probably due to some dynamical feedback effects.

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#### References

- Angell, J. K., Rocketsonde for a stratospheric temperature decrease in the western hemisphere during 1973–85, *Mon. Weather Rev.*, **115**, 2569–2577, 1987.
- Angell, J. K., Stratospheric warming due to Agung, El Chichon, and Pinatubo taking into account the quasi-biennial oscillation, *J. Geophys. Res.*, **102**, 9479–9485, 1997.
- Balachandra, N., and D. Rind, Modeling the effect of solar variability and the QBO on the troposphere/stratosphere system, I, The middle atmosphere, *J. Clim.*, **8**, 2058–2079, 1995.
- Chen, P., and W. A. Robinson, Propagation of planetary waves between the troposphere and stratosphere, *J. Atmos. Sci.*, **49**, 2533–2545, 1992.
- Dunkerton, T. J., and D. P. Delisi, Evolution of potential vorticity in the winter stratosphere of January–February 1979, *Geophys. Res. Lett.*, **91**, 1119–1208, 1986.
- Dunkerton, T. J., D. P. Delisi, and M. Baldwin, Middle atmosphere cooling trend in historical rocksonde data, *Geophys. Res. Lett.*, **25**, 3371–3374, 1998.
- Garcia, R., S. Solomon, R. G. Roble, and D. W. Rusch, A numerical response of the middle atmosphere to the 11-year solar cycle, *Planet. Space Sci.*, **32**, 411–423, 1984.
- Geller, M. A., and J. C. Alpert, Planetary wave coupling between the troposphere and the middle atmosphere as a possible sun-weather mechanism, *J. Atmos. Sci.*, **37**, 1197–1214, 1980.
- Harvey, V. L., and M. H. Hitchman, A climatology of the Aleutian High, *J. Atmos. Sci.*, **53**, 2088–2101, 1996.
- Hauchecorne, A., M. L. Chanin, and P. Keckhut, Climatology of the middle atmospheric temperature (30–90 km) and trends as seen by Rayleigh lidar above south of France, *J. Geophys. Res.*, **96**, 15 297–15 309, 1991.
- Hines, C. O., A possible mechanism for the production of sun-weather correlations, *J. Atmos. Sci.*, **31**, 589–591, 1974.
- Hood, L. L., J. L. Jirikovic, and J. P. McCormack, Quasi-decadal variability of the stratosphere: Influence of long term solar ultraviolet variations, *J. Atmos. Sci.*, **50**, 3941–3958, 1993.
- Hoxit, L. R. and Henry, R. M., Diurnal and annual temperature variations in the 30–60 km region as indicated by statistical analysis of rocketsonde temperature data, *J. Atmos. Sci.*, **30**, 922–933, 1973.
- Huang, T. Y. W., and G. P. Brasseur, Effect of long-term variability in a two-dimensional interactive model of the middle atmosphere, *J. Geophys. Res.*, **98**, 20 413–20 427, 1993.

- Karoly, D. J., and B. J. Hoskins**, Three-dimensional propagation of planetary waves, *J. meteor. Soc. Japan*, **60**, 109–122, 1982.
- Keckhut, P., A. Hauchecorne, and M. L. Chanin**, Midlatitude long-term variability of the middle atmosphere: trends and cyclic and episodic changes, *J. Geophys. Res.*, **100**, 18 887–18 897, 1995.
- Keckhut, P., M. E. Gelman, J. D. Wild, F. Tissot, A. J. Miller, A. Hauchecorne, M. L. Chanin, E. F. Fishbein, J. Gille, J. M. Russell III, and F. W. Taylor**, Semi-diurnal and diurnal temperature tides (30–55 km): climatology and effect on UARS-lidar data comparisons, *J. Geophys. Res.*, **101**, 10 299–10 310, 1996.
- Keckhut, P., F. J. Schmidlin, A. Hauchecorne, and M. L. Chanin**, Trend estimates from US rocketsondes at low latitude stations (8°S–34°N), taking into account instrumental changes and natural variability, submitted to *J. Atmos. Terr. Phys.*, **61**, 447–459, 1999.
- Kodera, K., and K. Yamazaki**, Long-term variations of upper stratospheric circulation in the northern hemisphere in December, *J. Meteorol. Soc. Jap.*, **68**, 101–105, 1990.
- Kodera, K., K. Yamazaki, M. Chiba, and K. Shibata**, Downward propagation of upper stratospheric mean zonal wind perturbation to the troposphere, *Geophys. Res. Lett.*, **17**, 1263–1266, 1990.
- Komuro, H.**, Long-term cooling in the stratosphere observed by aerological rockets at Ryori, Japan, *J. Meteorol. Soc. Jap.*, **68**, 1081–1082, 1989.
- Kokin, G. A., and E. V. Lysenko**, On temperature trends of the atmosphere from rocket and radiosonde data, *J. Atmos. Terr. Phys.*, **56**, 1035–1040, 1994.
- Labitzke, K., and M. P. McCormick**, Stratospheric temperature increases due to Pinatubo aerosols, *Geophys. Res. Lett.*, **19**, 207–210, 1992.
- McIntyre, M. E., and T. N. Palmer**, Breaking planetary waves in the stratosphere, *Nature*, **305**, 593–599, 1983.
- Miller, A. J., R. M. Nagatani, G. C. Tia, X. F. Niu, G. C. Reinsel, D. Wuebbles and K. Grant**, Comparisons of observed ozone and temperature trends in the lower stratosphere, *Geophys. Res. Lett.*, **19**, 929–932, 1992.
- Miller, A. J., S. M. Hollandsworth, L. E. Flynn, G. C. Tiao, G. C. Reinsel, L. Bishop, R. D. McPeters, W. G. Planet, J. J. DeLuisi, C. L. Mateer, D. Wuebbles, J. Kerr, and R. M. Nagatani**, Comparisons of observed ozone trends and solar effects in the stratosphere through examination of ground-based Umkehr and combined solar backscattered ultraviolet (SBUV) and SBUV2 satellite data, *J. Geophys. Res.*, **101**, 9017–9021, 1996.
- Mohanakumar, K.**, An investigation on the influence of solar cycle on mesospheric temperature, *Planet. Space Sci.*, **33**, 745–805, 1985.
- Newchurch, M. J., and D. M. Cunnold**, Aerosol effect on Umkehr ozone profiles using stratospheric aerosol and gas experiment II measurements, *J. Geophys. Res.*, **99**, 1383–1388, 1994.
- Pitari, G.**, A numerical study of the possible perturbation of stratospheric dynamics due to Pinatubo aerosols: implications for tracer transport, *J. Atmos. Sci.*, **50**, 2443–2461, 1993.
- O'Neill, A., W. L. Grose, V. D. Pope, H. Maclean, and R. Swinbank**, Evolution of the stratosphere during northern winter 1991/92 as diagnosed from U.K. meteorological office analyses, *J. Atmos. Sci.*, **51**, 2800–2817, 1994.
- Ramaswamy, V., M. D. Schwarzkopf, and J. Randel**, Fingerprint of ozone depletion in the spatial and temporal pattern of recent lower-stratospheric cooling, *Nature*, **382**, 616–618, 1996.
- Randel, W. J., J. C. Gille, A. E. Roche, J. B. Kumer, J. L. Mergenthaler, J. W. Waters, E. F. Fishbein, and W. A. Lahoz**, Stratospheric transport from the tropics to middle latitudes by planetary-wave mixing, *Nature*, **365**, 533–535, 1993.
- Randel, W. J., and J. B. Cobb**, Coherent variations of monthly mean total ozone and lower stratospheric temperature, *J. Geophys. Res.*, **99**, 5433–5447, 1994.
- Reinsel, G. C., G. C. Tiao, D. J. Wuebbles, J. B. Kerr, A. J. Miller, R. M. Nagatani, L. Bishop, and L. H. Ying**, Seasonal trend analysis of published ground-based and TOMS total ozone data through 1991, *J. Geophys. Res.*, **99**, 5449–5464, 1994.
- Rind, D., R. Suozzo, N. K. Balachandra, and M. J. Prather**, Climate change and the middle atmosphere. Part I: the impact of volcanic aerosols, *J. Clim.*, **5**, 189–207, 1992.
- Rind, D., N. K. Balachandra, and R. Suozzo**, Climate change and the middle atmosphere. Part II: the doubled CO<sub>2</sub> climate, *J. Atmos. Sci.*, **47**, 475–494, 1990.
- Sato, M., J. E. Hansen, M. P. McCormick, and J. B. Pollack**, Stratospheric aerosol optical depth, 1850–1990, *J. Geophys. Res.*, **98**, 22 987–22 994, 1993.
- Stratospheric processes and their role in climate/international ozone commission/global atmospheric watch (SPARC/IOC/GAW), Assessment of trends in the vertical distribution of ozone, SPARC Rep. 1, WMO Rep. 43, Eds. N. Harris, R. Hudson and C. Phillips, May 1998.
- Yata, A.**, Correction for temperature data obtained by MT-135 system, *Geophys. Mag.*, **35**, 99–121, 1970.
- World Meteorological Organization (WMO)**, Report of the international ozone trends panel 1988, global ozone research and monitoring project, 18, Geneva, 1988.
- World Meteorological Organization**, Scientific assessment of stratospheric ozone: 1989, global ozone research and monitoring project, 20, Geneva, 1989.